Electrical model of Cambrian rocks from Volodymyrska area in Volyno-Podillia (Ukraine)

This paper focuses on the techniques and results of electrical research into complex terrigenous and carbonate reservoirs. Presented here, are the electric data and their relation to the capacity properties of Cambrian sandstones, limestones and dolomites originating from the Volodymyrska area in Volyno-Podillia (Ukraine). Their petroelectrical models are generated.

Key words: reservoirs, porosity, permeability index, resistivity, formation resistivity factor, water saturation factor, resistivity increasing coefficient, sandstone, limestone and dolomite.

Introduction

The intricate correlation between capacity and filtration features, logging and field geophysical data requires a thorough analysis based on petrophysical laboratory study. A crucial factor in determining the geoelectrical properties of rocks is electrical resistivity ($\rho$), which is determined by rock composition and texture, capacity space structure, oil-gas- and water saturation of rocks, porosity factor, reservoir water salinity, temperature and pressure [1–8].

The main purpose was to develop an electrical model of reservoir rocks so as to provide the basis for a comprehensive analysis of their electrical parameters and their relation to capacity and filtration properties. Determining reservoir rock resistivity is essential for clarification of the variation range for certain types and groups of rocks, determining individual stratigraphic horizons, sections and facies; revealing the correlation between resistivity and a number of attributes, such as mineral composition, pore space structure, the phase relation of matter, frequency and tension of the electric field, as well as identifying the nature of changes in electrical resistivity under epigenetic transformations and metamorphic changes in rocks.

Laboratory data on resistivity variation in rocks are used in electrical logging interpretation and field electrical exploration.

This paper presents the results of petroelectrical laboratory analysis of Cambrian sandstones, limestones and dolomites from the Volodymyrska area, which is prospective for hydrocarbons (Volodymyrska-1 and Volodymyrska-2 wells, the interval 1190÷2520 m). The area is located in the northern part of the eastern side of the Lviv Paleozoic rock bend, in the Volyno–Podillia edge of the East-European platform.

The following summarizes petrographic characterization of typical species of the Cambrian rocks of the Volodymyrska area that is prospective for hydrocarbons.

1. Dolomites gray with a greenish tinge (intervals 1188÷1196 m, 1800÷1808 m and 1995÷2005 m) (Fig. 1a). The color is gray with a greenish tint. The texture is well equipped, poorly-layered. The rock is dense, slightly porous. The shear surface is flat and rough to the touch. The structure is finely crystalline. The composition of rocks: micrite dolomite, calcite, smectites.
2. Dolomitized limestone, organic-detritic, crinoidal (intervals 1820÷1824 m, 1831÷1842 m and 1885÷1895 m) (Fig. 1b). The laminated color changed from dark gray to brownish gray. The texture is foliated. Texture features are observed in polished sections. The shear surface is bumpy, with sharp edges, rough to the touch due to impurities of silt material. The rock is strong, slightly porous. The structure is cryptocrystalline (micrite), uniform in the areas of clusters of shells – organic-detritic. Ingredients: micrite dolomite, calcite, residues of calcite, there are few scales of chlorite and hydromica.

3. Silt sandstone, gray (interval 2189÷2200 m, 2210÷2219 m, 2263÷2276 m, 2296÷2304 m, 2520÷2526 m) (Fig. 1c). The color is gray; to a depth of 3 mm the color of the sample is ochre-brown. The texture is not foliated. The shear surface is rough, slightly bumpy, stepped. The rock is cemented, but dusty to the touch. The rock is porous, has the smell of hydrocarbons. The structure is fine-grained. The composition of rocks: debrises psammitic and silt dimension – 70%, cement is incomplete, micaceous-clayed, porous-opened.

Figure 1. The samples of rocks: a – dolomite, b – limestone, c – sandstone

**Experiment**

A series of laboratory experiments involved identifying the density of the rocks under study (dry and saturated with synthetic brine), open porosity (method of nitrogen saturation and method of synthetic brine saturation), residual water saturation factor (by centrifugation), permeability (nitrogen filtration method), interval time (velocity of P-waves) and resistivity. In the laboratory experiments, we determined electrical resistivity of rock samples under various conditions (dry, partially and completely saturated with reservoir synthetic brine) under atmospheric conditions and under those close to the reservoir.

Laboratory electrometric measurements of dry core samples were performed at a temperature of 20°C with a digital megohmmeter C.A. 6547, which ensures high-precision measurement of electrical resistivity in the range 10 kOhm to 10 TOhm, using a DC two-electrode scheme, with computerized digital recording [2, 3, 4, 5]. For NaCl (M = 30 g/l) saturated samples, RCL-meter MHC-1100 was used. Cylindrical samples to be tested were placed in a special core holder with nonpolarized electrodes, which are specially made from graphitized rubber. In order to determine the correlation between the petrophysical parameters and the water saturation levels (and hence oil and gas saturation) of rocks, we studied the changes in resistivity while stripping water on centrifuge OS-6M.

Electrical analysis involved repeated measuring of electrical resistivity in core samples saturated with synthetic brine. Measurements were performed before and after centrifugation in stripping modes from 1000 to 6000 rev/min. with a measurement pitch of 1000 rev/min., water displacement pressure ranging from 0.2 to 1.0 MPa (7 measurement cycles). Simultaneously, water saturation factor and the velocity of elastic waves were being determined. The mean relative error of electrical resistance was estimated to be 2.4% (permissible error is 5%).

**Data analysis.** Petrophysical laboratory research yielded data on porosity, permeability and electric properties of the major types of rocks, as shown in Table 1.

The laboratory measurements showed that resistivity values, measured on dry extracted samples (electrical resistivity of mineral skeleton) ranges from 0.2 to 1.0 MPa (7 measurement cycles). Simultaneously, water saturation factor and the velocity of elastic waves were being determined. The mean relative error of electrical resistance was estimated to be 2.4% (permissible error is 5%).
NaCl), ranges from 7.2 (sandstone) to 73 Ohm⋅m (limestone) with an average value of 45 Ohm⋅m.

The laboratory studies showed the following correlation (Fig. 2a, 2b, 2c) between porosity ($\phi$) and the formation resistivity factor ($FR$): $FR = a \cdot \phi^{-m}$ – Archie–Dahnov Equation for sandstone, limestone and dolomite, respectively, where $a$ is constant coefficient, and $m$ – structural indicator [2].

Electrical resistivity can be quantified using Archie–Dahnov Equation:

- For sandstone, $FR = 1.675 \cdot \phi^{-1.22}$ when $R^2 = 0.84$.
- For limestone, $FR = 40.99 \cdot \phi^{-0.27}$ when $R^2 = 0.77$.

Data analysis showed that in sandstones, changes in formation resistivity factor range from 20.4 to 85.5, with an average value of 50. Accordingly, porosity variation is within the range of 0.089 to 0.116, with an average value of 0.077. For limestone, variation in formation resistivity ranges from 143.9 to 207.6, with an average value of 180, whereas porosity varies from 0.002 to 0.015, with an average value of 0.006.

Table 1.

<table>
<thead>
<tr>
<th>Index #</th>
<th>Rock</th>
<th>Age</th>
<th>Parameter value</th>
<th>Density [kg/m³]</th>
<th>Open porosity [%]</th>
<th>Permeability $10^{-15}$·m²</th>
<th>Residual water saturation factor</th>
<th>Resistivity dry, NaCl</th>
<th>Resistivity saturated NaCl</th>
<th>Relative resistivity $F_R$</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>dry</td>
<td>saturated NaCl</td>
<td>by nitrogen</td>
<td>saturated NaCl</td>
<td>10⁻¹²·m²</td>
<td>dry, NaCl</td>
<td>saturated NaCl</td>
</tr>
<tr>
<td>1</td>
<td>limestone</td>
<td>min</td>
<td>2661</td>
<td>0.010</td>
<td>0.002</td>
<td>0.001</td>
<td>0.030</td>
<td>0.131</td>
<td>50.5</td>
<td>143.9</td>
</tr>
<tr>
<td>2</td>
<td>limestone</td>
<td>max</td>
<td>2699</td>
<td>0.025</td>
<td>0.013</td>
<td>3.479</td>
<td>0.670</td>
<td>6.588</td>
<td>72.9</td>
<td>207.6</td>
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<td>3</td>
<td>limestone</td>
<td>avg</td>
<td>2690</td>
<td>0.016</td>
<td>0.006</td>
<td>0.442</td>
<td>0.269</td>
<td>1.423</td>
<td>66.2</td>
<td>188.6</td>
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<tr>
<td>4</td>
<td>dolomite</td>
<td>min</td>
<td>2690</td>
<td>0.017</td>
<td>0.005</td>
<td>0.001</td>
<td>0.001</td>
<td>0.147</td>
<td>28.7</td>
<td>81.7</td>
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<td>5</td>
<td>dolomite</td>
<td>max</td>
<td>2845</td>
<td>0.046</td>
<td>0.038</td>
<td>0.002</td>
<td>0.650</td>
<td>24.442</td>
<td>69.8</td>
<td>198.7</td>
</tr>
<tr>
<td>6</td>
<td>dolomite</td>
<td>avg</td>
<td>2753</td>
<td>0.026</td>
<td>0.017</td>
<td>0.001</td>
<td>0.410</td>
<td>6.412</td>
<td>54.6</td>
<td>155.6</td>
</tr>
<tr>
<td>7</td>
<td>sandstone</td>
<td>min</td>
<td>2116</td>
<td>0.054</td>
<td>0.044</td>
<td>0.005</td>
<td>0.270</td>
<td>0.052</td>
<td>7.2</td>
<td>20.4</td>
</tr>
<tr>
<td>8</td>
<td>sandstone</td>
<td>max</td>
<td>2486</td>
<td>0.130</td>
<td>0.116</td>
<td>1.067</td>
<td>0.860</td>
<td>11.128</td>
<td>30.0</td>
<td>85.5</td>
</tr>
<tr>
<td>9</td>
<td>sandstone</td>
<td>avg</td>
<td>2290</td>
<td>0.089</td>
<td>0.077</td>
<td>0.139</td>
<td>0.710</td>
<td>3.648</td>
<td>15.5</td>
<td>44.1</td>
</tr>
</tbody>
</table>

Fig. 2. Correlation between porosity ($\phi$) and the formation resistivity factor ($F_R$) – Archie–Dahnov Equation (laboratory conditions): a – sandstone, b – limestone, c – dolomite.
of 0.008. Dolomites show variation in formation resistivity factor ranging from 81.7 to 198.7, with an average value of 145, and porosity variation of 0.005 to 0.038, with an average value of 0.017.

Limestone and dolomite from Volodymyrska area show little difference in the values: coefficient $a$ is 41 and 36.6 respectively, and the structural index $m = 0.273$ and 0.33. On the other hand, for sandstones from Volodymyrska area, coefficient $a$ is 1.675, and structural index $m = 1.124$, these values being markedly different from those above.

We used the centrifuge OS-6M in our laboratory experiment and carried out statistical analysis of electrical measurements, to determine the correlation dependences between resistivity index ($I_R$) and water saturation factor ($S_w$) for the rocks under study. Correlation dependences are as follows: for sandstones – $I_R = 1.069 \cdot S_w^{0.9972}$, with $R^2 = 0.82$, for the limestones the relationship is: $I_R = 1.169 \cdot S_w^{-0.79}$, with $R^2 = 0.86$; dolomites dependence can be expressed by the following formula: $I_R = 0.997 \cdot S_w^{0.87}$, with $R^2 = 0.79$, where $I_R = R/ R_o$, $R_t$ – partially water-saturated rock resistivity and $R_o$ is water-saturated rock resistivity. Figures 3a, 3b and 3c show the correlations.

Data analysis shows that the sandstones have a resistivity index ranging from 1 to 3.23, with an average value of 1.24. Accordingly, water saturation factor ranges from 1 to 0.29, with an average value of 0.88. For limestones, resistivity index variation is from 1 to 7.19, with an average value of 2.24 and water saturation factor ranging from 1 to 0.08, with an average value of 0.48. For dolomites, resistivity index ranges from 1 to 2.76, with an average value of 1.44 and the corresponding water saturation factor varying from 1 to 0.33, with an average value of 0.71.

It should be noted, that in the correlation equation $I_R = b \cdot S_w^n$ which expresses the relationship between water saturation factor and resistivity index, coefficient $b$ varies from 0.997 (dolomites) to 1.17 (limestones), and wettability index $n$ ranges from 0.79 (limestones) to 0.92 (sandstones). The rocks under study show small difference in these attributes.

To evaluate specific resistivity of rocks under reservoir conditions, we carried out a comprehensive study using a special installation for high pressure – VSC-1000, with pressure ranging from the atmospheric pressure to 59 MPa. The results clearly show that under increasing pressure, closing of micro-cracks and deformation of the pore space result in an increase of the electrical resistivity of rocks. We were able to define the correlation between mean resistivity increasing coefficient ($Q$) and pressure ($p$) for sandstones and limestones. This relationship can be expressed by polynomials of order 3 and 4:

For sandstones: $Q = 1 \cdot 10^{-5} \cdot p^3 - 13 \cdot 10^{-4} \cdot p^2 + 5.9 \cdot 10^{-2} \cdot p + 0.9993$, with $R^2 = 0.99$,

For limestones: $Q = -8 \cdot 10^{-6} \cdot p^4 + 9 \cdot 10^{-4} \cdot p^3 - 2.98 \cdot 10^{-2} \cdot p^2 + 4.022 \cdot p + 0.4302$, with $R^2 = 0.99$.

Figures 4a and 4b show this relationship.

**Fig. 3. Correlation dependence between water saturation factor ($S_w$) and resistivity index ($I_R$): a – sandstones, b – limestones, c – dolomites**
Sandstones show a variation of resistivity increasing coefficient ranging from 1 to 2.12, with an increase in the hydrostatic pressure from the atmospheric pressure up to 59 MPa. For the limestones, the variation range of resistivity increasing coefficient is from 1 to 8.3, with an increase in the hydrostatic pressure from the atmospheric pressure up to 59 MPa.

Graphs in Figure 4 suggest that there are three areas showing marked differences in the nature of electric resistivity variation. In the first area, where pressure ranges from the atmospheric pressure to 24.5 MPa, resistivity increasing coefficient is very high – up to 1.8 for sandstones. This may be caused by intense closing of microcracks, which reduces channels conductivity. The following range of pressure variation – from 24.5 to 44 MPa – is associated with a certain stabilization of electrical resistivity variation. In this case, the coefficient of resistivity increasing ranges from 1.8 to 2 for sandstones. In the range of pressures from 44 to 59 MPa resistivity increasing coefficient is lower than in the first range, but higher than in the second one. The slow growth rate of coefficient of resistivity increasing in the ranges (plots) 2 and 3 may be accounted for, by relatively smaller (than in the first range) deformations of the pore space, which constricts or breaks the conduction channels. For limestones, in the first section (pressure increases from atmospheric to 9.8 MPa) the coefficient of resistivity increasing varies from 1 to 2. In the range of pressures from 9.8 to 29 MPa, there is some stabilization in electrical resistivity variation to be observed. With pressure increasing above 30 MPa, the coefficient of resistivity increasing rises sharply (up to 8.3 at a pressure of 59 MPa). This is probably due to abrupt closing of major cracks responsible for electrical conductivity.

Using the data on measuring resistivity under pressure, we calculated its values for the rocks under reservoir conditions. A petroelectrical study with high pressures applied, enabled us to define the relationship between porosity (\(\phi\)) and formation resistivity factor \(F_R\). The Archie–Dahnov Equation for the Cambrian sandstones (under reservoir conditions) is as follows (Fig. 5): \(F_R = 1.365 \cdot \phi^{-1.52}\) with \(R^2 = 0.92\).

To evaluate electrical anisotropy, we took measurements of electrical resistivity along and across the stratification. Resistivity anisotropy coefficient \(\lambda\) was determined using the formula:

\[
\lambda = \frac{R_s}{R_v}
\]

where \(R_s\) and \(R_v\) – electrical resistivity along and across stratification respectively.

The results show that the resistivity anisotropy coefficient of dry extracted sandstones varies from 1.01 to 1.09 with an average value of 1.05. Anisotropy factor of saturated (solution of NaCl) sandstones varies from 1.05 to 1.18 with an average value of 1.12.

Comprehensive analysis of a petrophysical data set, was the basis for determining petroelectrical models of conser-
Table 2. Petroelectrical models of rocks (wells Volodymyrska-1 and Volodymyrska-2, the interval 1190–2520 m) from Volodymyrska area

<table>
<thead>
<tr>
<th>Index #</th>
<th>Petroelectrical parameter/Correlation dependence</th>
<th>Parameter variation range (average)/Correlation equation (correlation coefficient)</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Sandstone</td>
<td>Limestone</td>
</tr>
<tr>
<td>1</td>
<td>Electrical resistivity of dry extracted samples, MOhm · m</td>
<td>0.052 – 11.128 (3.647)</td>
</tr>
<tr>
<td>2</td>
<td>Electrical resistivity of rock samples, saturated with synthetic brine (solution of NaCl), MOhm · m</td>
<td>7.2 – 30 (17.6)</td>
</tr>
<tr>
<td>3</td>
<td>Formation resistivity factor ($F_b$) in laboratory conditions</td>
<td>20.4 – 85.8 (50)</td>
</tr>
<tr>
<td>4</td>
<td>Formation resistivity factor ($F_b$) under reservoir conditions</td>
<td>36.1 – 146.2 (75)</td>
</tr>
<tr>
<td>5</td>
<td>Archie–Dahnov dependence (laboratory conditions)</td>
<td>$F_b = 1.675 \cdot \phi^{-1.224}$ $R^2 = 0.84$</td>
</tr>
<tr>
<td>6</td>
<td>Archie–Dahnov dependence (under reservoir conditions)</td>
<td>$F_b = 1.365 \cdot \phi^{1.519}$ $R^2 = 0.92$</td>
</tr>
<tr>
<td>7</td>
<td>Resistivity anisotropy coefficient of dry extracted samples</td>
<td>1.01 – 1.09 (1.05)</td>
</tr>
<tr>
<td>8</td>
<td>Resistivity anisotropy coefficient of rock samples, saturated with synthetic brine (solution of NaCl)</td>
<td>1.05 – 1.18 (1.12)</td>
</tr>
<tr>
<td>9</td>
<td>Resistivity index ($I_b$)</td>
<td>1 – 3.23 (1.24)</td>
</tr>
<tr>
<td>10</td>
<td>Correlation between water saturation factor ($S_w$) and resistivity index ($I_b$)</td>
<td>$I_b = 1.069 \cdot S_w^{-0.82}$ $R^2 = 0.82$</td>
</tr>
<tr>
<td>11</td>
<td>Correlation between the of resistivity increasing coefficient ($Q$) and pressure ($p$)</td>
<td>$Q = 1 \cdot 10^{5} \cdot p^3 + 13 \cdot 10^4 \cdot p^2 + 5.9 \cdot 10^2 \cdot p + 0.9993$ $R^2 = 0.99$</td>
</tr>
<tr>
<td>12</td>
<td>Resistivity increasing coefficient with pressure ranging from atmospheric pressure to 59 MPa</td>
<td>1 – 2.12</td>
</tr>
</tbody>
</table>

Conclusions. It was shown that, being a powerful tool for both laboratory and field studies, geoelectric methods prove to be effective and provide extensive and accurate data on the properties of rocks. Electrical study plays an important role in petrophysics and is widely used in determining the physical properties of rocks. It is obvious, that for solving various tasks in the search and exploration of mineral deposits, particularly in petroleum geology, it is essential in determining the composition, structure and condition of rocks.

Laboratory experiment was conducted in order to determine the electrical attributes of the Cambrian sandstones, limestones and dolomites from Volodymyrska area, which is prospective for hydrocarbons. Our study has revealed the empirical relationships between electrical parameters and filtration–capacitive properties of sandstones, limestones and dolomites, which are essential for the geological interpretation of geophysical data. These relationships are approximated by a power function.

Detailed electrical research into the properties of sandstones, limestones and dolomites has ensured accurate electrical models of these rocks based on geological and geophysical data. The models show significant differences in the electrical parameters of sandstones, limestones and dolomites. These models can become a powerful tool in studying the physical properties of different types of rocks. Further research into electrical properties of rocks will require data on dielectric permeability, dielectric loss tangent, as well as evaluating...
the relevance of of geoelectric parameters which account for variation in electrical resistivity of dry, extracted samples exposed to direct current over longer periods of time and establishing their correlations with logging data.

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Literature


Serhiy VYZHVA
Professor, Doctor of Geology
Head of the Department of Geophysics
Distinguished Educationalist of Ukraine
ESI “Institute of Geology”, Taras Shevchenko National University of Kyiv
Ukraine, 03022, Kyiv, 90 Vasylkivska str.
E-mail: vyzhva_s@ukr.net

Viktor ONYSHCHUK
Ph.D. in Geology
Assistant of the Department of Geoinformatics
ESI “Institute of Geology”, Taras Shevchenko National University of Kyiv
Ukraine, 03022, Kyiv, 90 Vasylkivska str.
E-mail: vitus16@ukr.net

Dmytro ONYSHCHUK
Ph.D. in Geology
Standardization, certification and quality specialist LLC “Ancor Personnel Ukraine”
Ukraine 03680, Kyiv, Business Center Olimpiyskyi, 72 Velyka Vasylkivska str.
E-mail: dmytrij48@gmail.com